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# The global N<sub>2</sub>O Model Intercomparison Project (NMIP): Objectives, Simulation Protocol and Expected Products

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43 **Abstract**

44 Nitrous oxide (N<sub>2</sub>O) is an important greenhouse gas (GHG) and also an ozone-depleting  
45 substance that has both natural and anthropogenic sources. Large uncertainty remains on the  
46 magnitude and spatiotemporal patterns of N<sub>2</sub>O fluxes and the key drivers of N<sub>2</sub>O production in  
47 the terrestrial biosphere. Some terrestrial biosphere models have been evolved to account for  
48 nitrogen processes and show the capability to simulate N<sub>2</sub>O emissions from land ecosystems at  
49 the global scale, but large discrepancies exist among their estimates primarily due to inconsistent  
50 input data sets, simulation protocol, and model structure and parameterization schemes. Based on  
51 the consistent model input data and simulation protocol, the global N<sub>2</sub>O Model Inter-Comparison  
52 Project (NMIP) was initialized with ten state-of-the-art terrestrial biosphere models with N  
53 cycling included. Specific objectives of NMIP are to: 1) Unravel the major N cycling processes  
54 controlling N<sub>2</sub>O fluxes in each model and identify the uncertainty sources from model structure,  
55 input data and parameters; 2) Quantify the magnitude, spatial and temporal patterns of global and  
56 regional N<sub>2</sub>O fluxes from the pre-industrial period (1860) to present, and attribute the relative  
57 contributions of multiple environmental factors to N<sub>2</sub>O dynamics; and 3) Provide a bench-  
58 marking estimate of N<sub>2</sub>O fluxes through synthesizing the multi-model simulation results and  
59 existing estimates from ground-based observations, inventories, and statistical and empirical  
60 extrapolations. This study provides detailed descriptions for the NMIP protocol, input data,  
61 model structure and key parameters, along with preliminary simulation results. The global and  
62 regional N<sub>2</sub>O estimation derived from the NMIP is a key component of the Global N<sub>2</sub>O Budget  
63 activity jointly led by the Global Carbon Project (GCP) and the International Nitrogen Initiative  
64 (INI).

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Capsules

The N<sub>2</sub>O Model Inter-Comparison Project (NMIP) aims at understanding and quantifying the budgets of global and regional terrestrial N<sub>2</sub>O fluxes, environmental controls and uncertainties associated with input data, model structure and parameters.

## 89 **1. Introduction**

90 Nitrous oxide (N<sub>2</sub>O) is an important greenhouse gas and the time-integrated radiative  
91 forcing resulting from a mass unit of N<sub>2</sub>O is 298 times larger than that from carbon dioxide (CO<sub>2</sub>)  
92 emissions for a 100-year time horizon (Ciais et al., 2013; Myhre et al., 2013). Multiple lines of  
93 evidence indicate that human activities (e.g., industrial N<sub>2</sub> fixation by the Haber-Bosch process  
94 or by combustion, and manure N application) play an increasingly significant role in the  
95 perturbation of the global N cycle (Galloway et al., 2008; Gruber et al., 2008; Fowler et al.,  
96 2015), which has led to an increase in atmospheric N<sub>2</sub>O concentration by ~21%, from 271 ppb at  
97 pre-industrial level to 329 ppb in 2015 (MacFarling et al., 2006; Prather et al., 2012, 2015;  
98 Thompson et al., 2014; <https://www.esrl.noaa.gov/>). The anthropogenic N<sub>2</sub>O emissions are  
99 estimated to have increased from 0.7 Tg N yr<sup>-1</sup> in 1860 to 6.9 Tg N yr<sup>-1</sup> in 2006, ~60% of which  
100 was ascribed to agricultural activities (Ciais et al., 2013, Davidson and Kanter, 2014). The  
101 increased N<sub>2</sub>O emissions have significantly contributed to climate warming. During the 2000s,  
102 the warming effect of N<sub>2</sub>O emissions from the terrestrial biosphere counteracted more than half  
103 of the cooling effect of the global land CO<sub>2</sub> sink (Tian et al., 2016), and anthropogenic N<sub>2</sub>O  
104 emissions are projected to lead to further global warming during the 21<sup>st</sup> century and beyond  
105 (Stocker et al., 2013).

106 In terrestrial ecosystems, N<sub>2</sub>O is mainly produced in soils via nitrification and  
107 denitrification processes (Smith and Arah, 1990; Wrage et al., 2001; Schmidt et al., 2004). All  
108 these processes are regulated by microbial activities under various soil micro-environments such  
109 as soil temperature, moisture and aeration, clay content, pH, and C and N availability (Firestone  
110 and Davidson, 1989; Goldberg and Gebauer, 2009; Butterbach-Bahl et al., 2013; Brotto et al.,  
111 2015; Rowlings et al., 2015). In addition, N<sub>2</sub>O emissions from terrestrial ecosystems can be

112 regulated by both natural disturbances and human management such as synthetic N fertilizer,  
113 manure N application, irrigation, tillage, and the choice of crop varieties (Rice and Smith, 1982;  
114 Cai et al., 1997; Ding et al., 2010). However, our understanding of the mechanisms responsible  
115 for terrestrial N<sub>2</sub>O emissions is still limited, which contributes to large uncertainties in estimating  
116 both preindustrial and contemporary N<sub>2</sub>O emissions. For example, estimates of global terrestrial  
117 N<sub>2</sub>O emissions from natural sources vary by up to a factor of three and range between 3.3 and  
118 9.0 Tg N yr<sup>-1</sup> (Ciais et al. 2013). Human-induced biogenic N<sub>2</sub>O emissions from the land  
119 biosphere have not yet been investigated well (Tian et al. 2016). Therefore, a major international  
120 and multidisciplinary effort is required to assess information from different research disciplines  
121 and approaches in order to constrain current knowledge on the N<sub>2</sub>O budget and drivers, and to  
122 identify research gaps.

123         Process-based modeling is an essential tool in assessing and predicting the terrestrial N  
124 cycle and N<sub>2</sub>O fluxes in response to multi-factor global changes. Several process-based models  
125 have been used to estimate N<sub>2</sub>O emissions from natural and agricultural soils at various  
126 spatiotemporal scales. The conceptual model of “hole-in-pipe” (Firestone and Davidson, 1989)  
127 was first incorporated in the Carnegie-Ames-Stanford-Approach (CASA) Biosphere model  
128 (Potter et al., 1993) to estimate N trace gas emissions at the global scale (Potter et al., 1996). The  
129 daily version of the CENTURY model (DAYCENT) was linked to atmospheric models to better  
130 estimate trace gas fluxes from different ecosystems (Parton et al., 1998). The DeNitrification-  
131 DeComposition model (DNDC; Li et al., 1992) was developed to study the impacts of various  
132 agricultural practices on N<sub>2</sub>O emissions. In the Dynamic Land Ecosystem Model (DLEM), Tian  
133 et al. (2011, 2015) considered the biotic and abiotic processes (e.g., plant N uptake and N  
134 leaching loss) that regulate N<sub>2</sub>O fluxes in natural and managed soils. In recent years, multiple C-

135 N coupled models, such as DyN-LPJ (Xu-Ri and Prentice, 2008), O-CN (Zaehle and Friend,  
136 2010, 2011), Land Surface Processes and exchanges model of the University of Bern (LPX-Bern  
137 1.0; Stocker et al., 2013), CLMCN-N<sub>2</sub>O (Saikawa et al., 2014), and LM3V-N (Huang and Gerber,  
138 2015) were developed by integrating a prognostic N cycle into different land surface models and  
139 simulate N<sub>2</sub>O emissions from land ecosystems. Unsurprisingly, these models generated divergent  
140 estimates of global terrestrial N<sub>2</sub>O budgets and spatiotemporal patterns mainly due to differences  
141 in model input datasets, model structure, and parameterization schemes. What are the major  
142 contributing factors responsible for the changing patterns of terrestrial N<sub>2</sub>O emissions? How can  
143 we narrow down the model-estimated bias or uncertainties? What are the knowledge gaps in  
144 fully accounting for the N<sub>2</sub>O processes? Here, we attempt to answer these questions through the  
145 establishment and designing of the global N<sub>2</sub>O Model Inter-comparison Project (NMIP).

146         During the past two decades, carbon-related Model Intercomparison Projects (MIPs) have  
147 been established to evaluate model uncertainties in simulating the terrestrial carbon dynamics.  
148 For example, the Vegetation-Ecosystem Modeling and Analysis Project (VEMAP) was a pioneer  
149 MIP activity, driven by a common model input database and was established to provide multi-  
150 model ensemble estimates of carbon fluxes and storage in response to changing climate and  
151 atmospheric CO<sub>2</sub> (Melillo et al., 1995; Schimel et al., 2000). More recently, a number of CO<sub>2</sub>-  
152 oriented MIPs and synthesis activities were implemented, such as the North American Carbon  
153 Program site and regional synthesis (NACP; Schwalm et al., 2010; Richardson et al., 2011;  
154 Schaefer et al., 2012) and its extended Multi-Scale Synthesis and Terrestrial Model  
155 Intercomparison Project (MsTMIP; Huntzinger et al., 2013; Wei et al., 2014), the TRENDY  
156 Project (Le Quéré et al., 2016; Sitch et al., 2015), the Inter-Sectoral Impact Model  
157 Intercomparison Project (ISI-MIP) (Warszawski et al., 2014; Ito et al., 2016), and the Multi

158 Model-data Synthesis of Terrestrial Carbon Cycles in Asia (Asia-MIP; Ichii et al., 2013). These  
159 MIPs enhanced our understanding of model uncertainties and provided insight into future  
160 directions of model improvement.

161         Following the CO<sub>2</sub>-related MIPs, global methane (CH<sub>4</sub>) MIPs and synthesis activities  
162 were implemented in recent years, for example, the Wetland and Wetland CH<sub>4</sub> Intercomparison  
163 of Models Project (WETCHIMP; Melton et al., 2013; Wania et al., 2013) and GCP global CH<sub>4</sub>  
164 budget synthesis (Sauniois et al., 2016; Poulter et al. 2017). Although terrestrial biogenic N<sub>2</sub>O  
165 emissions significantly contribute to climate warming, the model development for simulating N  
166 cycle and N<sub>2</sub>O fluxes remains far behind the CO<sub>2</sub>- and CH<sub>4</sub>-related activities. The relatively  
167 sparse and short-term observations limited our understanding of N cycling in terrestrial  
168 ecosystems. Comparing with CO<sub>2</sub> and CH<sub>4</sub>, smaller N<sub>2</sub>O concentration gradients in the  
169 atmosphere and the varying magnitudes of soil N<sub>2</sub>O emissions across observation sites and  
170 periods makes it more difficult to quantify the N<sub>2</sub>O budget at a large scale. Another important  
171 uncertainty comes from the differences in model representation and parameterization schemes of  
172 N processes, and the influence of biophysical and environmental factors on N<sub>2</sub>O dynamics  
173 (Appendix A). Similar to the purposes of the CO<sub>2</sub>- and CH<sub>4</sub>-related MIPs, there is a need to  
174 initialize a MIP for the N models to assess the global N<sub>2</sub>O budget. Under the umbrella of the  
175 Global Carbon Project (GCP) and the International Nitrogen Initiative (INI), we initiated the  
176 NMIP to investigate the uncertainty sources in N<sub>2</sub>O estimates and provide multi-model N<sub>2</sub>O  
177 emissions estimates from natural and agricultural soils. This paper describes the detailed NMIP  
178 protocol, input data, model structure, and some preliminary simulation results.

179

## 180 **2. The NMIP framework, objectives and tasks**



181 Motivated by large uncertainties and increasing data availability, the NMIP is developed  
182 to establish a research network for providing a multi-model ensemble estimate on the  
183 global/regional N<sub>2</sub>O budgets and to identify major uncertainties associated with model structure,  
184 parameters and input data (Figure 1). This project was first proposed at the Regional Carbon  
185 Cycle Assessment and Processes (RECCAP) workshop “4<sup>th</sup> International workshop on Asian  
186 Greenhouse Gases”, in JAMSTEC, Yokohama, Japan, April 8-10, 2014. The NMIP was  
187 launched at a side meeting during the 2015 American Geophysical Union fall meeting and began  
188 to work in the fall of 2016.

189 Specific objectives of NMIP are: 1) Unravel the major N cycling processes controlling  
190 N<sub>2</sub>O fluxes in each model and identify the uncertainty sources from modeling structure, input  
191 data and parameters; 2) Quantify the magnitude, spatial and temporal patterns of global and  
192 regional N<sub>2</sub>O fluxes during 1860–2015, and attribute the relative contributions of multiple  
193 environmental factors to N<sub>2</sub>O dynamics; and 3) Provide a bench-mark estimate of global/regional  
194 N<sub>2</sub>O fluxes through synthesizing the multi-model simulation results and existing estimates from  
195 ground-based observations, inventories, and statistical/empirical extrapolations. To achieve these  
196 objectives, the NMIP group members have collectively developed a model simulation protocol  
197 as outlined in Figure 1.

198 There are five key tasks or progressing stages in the protocol: 1) Development and  
199 delivery of spatiotemporal model driving forces; 2) Individual model calibration and evaluation;  
200 3) Model simulations and delivery of results; 4) Quality control and analysis of model results; 5)  
201 Synthesis and uncertainty analysis.

202

### 203 **3. Key model input datasets**

204 To minimize the uncertainty that results from input datasets, the NMIP provided  
205 consistent model driving datasets for all modeling groups. The datasets include potential  
206 vegetation, climate, atmospheric CO<sub>2</sub> concentration, atmospheric N deposition, synthetic N  
207 fertilizer applications in cropland and pasture, manure N production and applications in cropland  
208 and pasture, and historical distribution of cropland at a spatial resolution of 0.5° by 0.5°  
209 latitude/longitude (Table 1). Half-degree resolution is appropriate for studies at a global scale,  
210 considering that most of model input data are available and many previous MIPs at a global scale  
211 were conducted at this resolution. Here we briefly describe these input data sets and their  
212 sources.

213 1) Climate: CRU-NCEP climate version 7 is a fusion of the CRU and NCEP/NCAR  
214 reanalysis climate datasets between 1901 and 2015, which was reconstructed by the Laboratoire  
215 des Sciences du Climat et l'Environnement, Paris, France (<https://vesg.ipsl.upmc.fr>). Major  
216 climate variables include longwave and shortwave radiation, air pressure, humidity, temperature,  
217 precipitation, and wind speed at 6-hourly temporal resolution. Monthly magnitude of climate  
218 variables in CRU-NCEP dataset was forced to be consistent with the observational-based CRU  
219 datasets.

220 2) Atmospheric CO<sub>2</sub>: Monthly atmospheric CO<sub>2</sub> concentration from 1860 to 2015 was  
221 obtained from NOAA GLOBALVIEW-CO<sub>2</sub> dataset derived from atmospheric and ice core  
222 measurements (<https://www.esrl.noaa.gov>).

223 3) Vegetation: Potential vegetation map was acquired from the Synergetic Land Cover  
224 Product (SYNMAP, <ftp://ftp.bgc-jena.mpg.de/pub/outgoing/mjung/SYNMAP/>), which merged  
225 multiple satellite-global land cover maps into a desired classification approach (Jung et al., 2006).  
226 Each 0.5° grid cell includes the area fractions for a maximum of 47 land cover types. Vegetation

227 in SYNMAP is classified according to its life form, leaf type, and leaf longevity. Barren ground,  
228 permanent snow and ice are also included in this dataset. Based on this SYNMAP dataset,  
229 participating model groups could create vegetated land fraction and reorganize the vegetation  
230 types to generate the corresponding plant functional type and fractions for their models. Annual  
231 cropland area from 1860 to 2015 was acquired from the HYDE 3.2 datasets  
232 (<ftp://ftp.pbl.nl/hyde/>), which reconstructed time-dependent land use by historical population and  
233 allocation algorithms with weighting maps (Klein Goldewijk et al., 2016). This dataset shows  
234 that global cropland area increased from 5.9 million km<sup>2</sup> in the year of 1850 to 15.2 million km<sup>2</sup>  
235 in the year of 2015.

236 4) Atmospheric N deposition onto land surface: The monthly atmospheric N deposition  
237 (NH<sub>x</sub>-N and NO<sub>y</sub>-N) during 1860 - 2014 were from the IGAC/SPARC Chemistry-Climate  
238 Model Initiative (CCMI) N deposition fields. CCMI models explicitly considered N emissions  
239 from natural biogenic sources, lightning, anthropogenic and biofuel sources, and biomass  
240 burning (Eyring et al., 2013). The transport of N gases has also been simulated by the chemical  
241 transport module in CCMI models. This data was recommended by the Coupled Model  
242 Intercomparison Project (CMIP) and used as the official products for CMIP6 models that lack  
243 interactive chemistry components ([https://blogs.reading.ac.uk/ccmi/forcing-databases-in-support-](https://blogs.reading.ac.uk/ccmi/forcing-databases-in-support-of-cmip6/)  
244 [of-cmip6/](https://blogs.reading.ac.uk/ccmi/forcing-databases-in-support-of-cmip6/)).

245 5) N fertilizer application: Spatially-explicit synthetic N fertilizer use data was  
246 specifically developed in this project. We reconstructed the annual synthetic/mineral N fertilizer  
247 dataset from 1960 to 2014 for the global cropland, matched with HYDE 3.2 cropland distribution  
248 (Lu and Tian, 2017; <https://doi.pangaea.de/10.1594/PANGAEA.863323>). Data on national-level  
249 crop-specific fertilizer use amount was collected from the International Fertilizer Industry

250 Association (IFA) and FAO. This N fertilizer dataset shows that the global total N fertilizer  
251 amount increased from 11 Tg N yr<sup>-1</sup> in 1960 to 112 Tg N yr<sup>-1</sup> in 2013, and N fertilizer use rate  
252 per unit cropland area increased by about 8 times in this period. N fertilizer application rate  
253 before 1960 was linearly reduced to the zero in the 1900s.

254 6) Manure N production and application: Gridded annual manure N production in the  
255 period of 1860-2014 was developed by integrating the Global Livestock Impact Mapping System  
256 (GLIMS), the country-level livestock population from FAO, and N excretion rates of different  
257 livestock categories according to IPCC 2006-Tier I and (Zhang et al., 2017,  
258 <https://doi.org/10.1594/PANGAEA.871980>). This annual dataset shows that manure N  
259 production increased by more than 6 times from 21 Tg N yr<sup>-1</sup> in 1860 to 131 Tg N yr<sup>-1</sup> in 2014,  
260 and application rate of manure N to cropland is less than 20% of the total production. In this  
261 project, we only consider the manure N application in cropland area. Manure N production and  
262 application rates in 2015 was assumed to be same as that in 2014.

263 All the input datasets were delivered to the modeling groups in netCDF format. To fit  
264 with individual modeling requirements for input datasets, the modeling groups could either use a  
265 subset of these data sets or add some additional data sets. For example, the participating model  
266 DLEM used all these environmental factors as inputs, while the model O-CN did not use manure  
267 N as an input (See Table 3 for model input requirements in each model). Figure 2 illustrates the  
268 inter-annual variations of the major input datasets at the global level during different available  
269 time periods. Figure 3 shows the spatial patterns of atmospheric N deposition, N fertilizer use,  
270 and manure N production in 1860, 1900, 1950, and 2014.

271

#### 272 **4. Model result benchmarking and evaluation**

273           Except for bottom-up model simulations, the NMIP also plans to synthesize multiple  
274 sources of terrestrial soil N<sub>2</sub>O emission data to provide a benchmark for evaluating model  
275 estimates. Four types of data will be collected or developed to serve as a potential benchmark: 1)  
276 Site-level N cycling processes and N<sub>2</sub>O emission measurements through chamber or eddy-flux  
277 tower across biomes; 2) N<sub>2</sub>O flux measurement data from national-based or global-based  
278 measurement network (e.g., Long Term Ecological Research Network, Long Term  
279 Agroecosystem Research Network, Greenhouse gas Reduction through Agriculture Carbon  
280 Enhancement network, <http://www.n2o.net.au>, etc.) ; 3) Other spatialized datasets, including  
281 statistical extrapolation (e.g., Xu et al. 2008; Kurokawa et al., 2013; Zhuang et al., 2012); 4) N<sub>2</sub>O  
282 fluxes from other than terrestrial ecosystem sources to allow for a global budget (industrial,  
283 combustion, waste water & water bodies, marine and oceanic sources) (e.g., Battaglia and Joos,  
284 in press; Davidson and Kanter, 2014; Galloway et al., 2004; Fowler et al., 2013; Winiwarter et  
285 al., 2017); and 5) atmospheric inversions (e.g., Saikawa et al., 2014; Thompson et al., 2014) in  
286 conjunction with atmospheric N<sub>2</sub>O measurements from tall towers. We also call for more  
287 observation-derived studies to provide regional and global N<sub>2</sub>O emission estimates through  
288 advanced computational techniques, such as machine learning, Multi-Tree Ensemble (MTE), and  
289 remote sensing products. We anticipate that through multiple constraints, process-based  
290 modeling approach can be more effective and reliable in estimating magnitude, spatial and  
291 temporal patterns of terrestrial N<sub>2</sub>O emissions, and quantifying relative contributions of  
292 environmental drivers to N<sub>2</sub>O dynamics.

293

## 294 **5. Major characteristics of participating models**

295           The N cycle in the earth system involves complex biogeochemical processes, in which N  
296 is transformed into various chemical forms, and circulates among the atmosphere, terrestrial and  
297 aquatic ecosystems. Important terrestrial processes in the N cycle include biological N fixation  
298 (BNF), mineralization (conversion of organic N to inorganic N during the processes of organic  
299 matter decomposition), immobilization (transformation of soil inorganic N to organic N),  
300 ammonification (conversion of organic to ammonium N), volatilization (transformation of soil  
301 ammonium N to ammonia gas), nitrification (transformation of ammonium N to nitrate and  
302 nitrite N), denitrification (the process of nitrate reduction by microbial activities), plant uptake  
303 from soil, resorption by living plant organs, adsorption and desorption by soil mineral particles,  
304 and N leaching from soil to aquatic systems. The modeled N processes include N transformation  
305 between organic and inorganic forms and movements among atmosphere, vegetation, soil, and  
306 riverine systems. Although N processes are tightly coupled with carbon processes in soil and  
307 vegetation, the greater variability in N processes compared to C processes make it more difficult  
308 to simulate N cycling. At current stage, the NMIP has included ten ecosystem models with  
309 explicit terrestrial N cycling processes (Table 2; Figure 1). Nine models (DLEM, LM3V-N,  
310 ORCHIDEE, ORCHIDEE-CNP, O-CN, LPJ-GUESS, LPX-Bern, TRIPLEX-GHG and VISIT)  
311 are capable of simulating N<sub>2</sub>O emissions from both natural and agriculture ecosystems, while  
312 one model (CLM-CN) only simulates N<sub>2</sub>O emissions from natural vegetation. The biophysical  
313 processes (such as, canopy structure, albedo and evapotranspiration), biogeochemical processes  
314 (such as decomposition and denitrification), and N input for cropland are significantly different  
315 from those for natural vegetation. For example, temperature in cropland was found to be lower  
316 than that in natural forest due to the higher albedo and evapotranspiration (Bonan, 2000). These  
317 differences could lead to different magnitude and timing of N<sub>2</sub>O emissions from cropland.

318 Therefore, biophysical characteristics and management practices in cropland, such as crop  
319 cultivation, fertilizer uses, irrigation and harvesting, are required to be explicitly represented by  
320 the models with crop module.

321 In order to assess the uncertainty from model structure, each participating model was  
322 asked to complete a detailed survey specifying the modeling mechanisms in exogenous N inputs  
323 (e.g., N deposition, synthetic N fertilizer and manure N application, and BNF) and N  
324 transformation processes. The summarized survey results are shown in Table 3. In general, N<sub>2</sub>O  
325 emissions from soil are regulated at two levels, which are the rates of nitrification and  
326 denitrification in the soil and soil physical factors regulating the ratio of N<sub>2</sub>O to other nitrous  
327 gases (Davidson et al., 2000).

328 For N input to land ecosystems, all ten models considered the atmospheric N deposition  
329 and biological fixation, nine models with a crop N<sub>2</sub>O module included N fertilizer use, but only  
330 six models considered manure as N input. For vegetation processes, all models included dynamic  
331 algorithms in simulating N allocation to different living tissues and vegetation N turnover, and  
332 simulated plant N uptake using the “Demand and Supply-driven” approach. For soil N processes,  
333 all ten models simulated N leaching according to water runoff rate; however, the models differ in  
334 representing nitrification and denitrification processes and the impacts of soil chemical and  
335 physical factors. The differences in simulating nitrification and denitrification processes are one  
336 of the major uncertainties in estimating N<sub>2</sub>O emissions. Algorithms associated with N<sub>2</sub>O  
337 emissions in each participating model are briefly described in Appendix A.

338

## 339 **6. The NMIP model simulation methods and experimental designs**

### 340 **6.1 Model initialization**

341 The model simulations were divided into two stages: (1) spin-up and (2) transient runs  
342 (Figure 4). During the spin-up run, models were driven by the repeated climate data from 1901 -  
343 1920 and by other driving forces in 1860, i.e., atmospheric CO<sub>2</sub> concentration, N deposition, N  
344 fertilizer use, manure N application, and land cover and land use change (LCLU). The N  
345 fertilizer use was assumed to be zero in 1860. Each model group could determine the spin-up  
346 running years according to model's specific requirement. For example, the DLEM assumed that  
347 model reaches the equilibrium status when the differences of grid-level carbon, N, and water  
348 stocks were less than 0.5 g C m<sup>-2</sup>, 0.5 g N m<sup>-2</sup>, and 0.5 mm in two consecutive 50 years. When  
349 these thresholds were met, the spin-up run stopped and model reached an equilibrium state.

350

## 351 **6.2 Model simulation experiments**

352 During the transient run, seven experiments were designed to simulate global terrestrial  
353 N<sub>2</sub>O emissions. All the model experiments started with the equilibrium carbon, water and N  
354 status in 1860, which is obtained from the spin-up run, and transiently ran through the period  
355 during 1860-2015 (Figure 4). For the period of 1860-1900 when CRU/NCEP climate data is not  
356 available, the 20-year average climate data between 1901 and 1920 were used. In the NMIP, we  
357 applied the progressively reducing factor experimental scheme (i.e., first experiment includes all  
358 factors and then reduce one factor each time; the effect of this factor is equal to the difference  
359 between previous and current experiment) to simulate the impacts of individual environmental  
360 factors on N<sub>2</sub>O fluxes. In total, seven experiments (from S0 to S6) were designed (Figure 4). The  
361 S0 reference (baseline) run was designed to track the model internal fluctuation and model drift.  
362 The S1 experiment included the temporal variations of all time-varying driving forces. “Best  
363 estimate” of N<sub>2</sub>O emissions were acquired from either S1 experiment (for models considering



364 manure as input) or S2 experiment (for models without considering manure). The overall effect  
365 of all environmental factors was calculated as: S1-S0. The effects of manure N use (MANN), N  
366 fertilizer use (NFER), N deposition (NDEP), LCLU, atmospheric CO<sub>2</sub> (CO<sub>2</sub>), and climate (CLIM)  
367 were respectively calculated as: S1-S2, S2-S3, S3-S4, S4-S5, S5-S6, and S6-S0, respectively.

368

## 369 **7. Model outputs, quality control and data availability**

370 All participating model groups are requested to provide the gridded simulations of N<sub>2</sub>O  
371 fluxes from global terrestrial ecosystems and other relevant variables that can be used for  
372 understanding C-N coupling and key N processes simulated by each individual model (Table 4).  
373 The required model output will be submitted at the annual time-step during 1860-2015 and at the  
374 monthly time-step during 1980-2015. In addition to modeling estimates of grid-level fluxes and  
375 pool sizes, modeling groups will submit biome-level results to facilitate biome-level N<sub>2</sub>O  
376 emission analysis and split contributions of global N<sub>2</sub>O dynamics to primary biome types. The  
377 model output from each modeling group is sent to the core team led by Dr. Hanqin Tian for data  
378 quality checking and preliminary analysis. The quality control is conducted to check if the  
379 individual model results are reasonable and to avoid the obvious errors during model simulations.  
380 After quality control process, model output is transferred to a data sharing website.

381 The model input and output datasets are made available to all model groups for further  
382 analyses. Model input data and model results will be made available to broader research  
383 community once the results of the first NMIP are published. A data use and authorship policy  
384 has been established.

385

## 386 **8. Result analysis and synthesis**

387           Based on model results, the NMIP team will provide multi-model ensemble estimates for  
388 terrestrial N<sub>2</sub>O fluxes at various scales from country-, sector-, continental, to global, and also  
389 assess differences and uncertainties among participating models. Through the seven simulation  
390 experiments, the magnitudes and spatiotemporal variations in terrestrial N<sub>2</sub>O emissions will be  
391 attributed to changes in different environmental factors at both regional and global scales. The  
392 global and regional N<sub>2</sub>O flux data derived from other sources including atmospheric inversion,  
393 statistical extrapolation, and inventory approaches (e.g., the N<sub>2</sub>O emission data collected in Tian  
394 et al. 2016) will be compared and integrated with the NMIP modeled results. Through these  
395 syntheses and evaluations of modeled versus field observed N<sub>2</sub>O dynamics, we will further  
396 identify the gaps in our understanding to estimate N<sub>2</sub>O fluxes and put forward potential strategies  
397 to improve the models. In the following sections, we provide an initial analysis of simulated  
398 terrestrial N<sub>2</sub>O emissions from the three models (DLEM, O-CN, and VISIT), which simulate  
399 both natural and agricultural emissions.

400           As indicated by the model ensemble, the global N<sub>2</sub>O emission has significantly increased,  
401 especially since the 1960s with more rapidly rising exogenous N inputs to terrestrial ecosystems  
402 (Figure 5). Natural soils were the largest source across the entire period. Cropland is the single  
403 largest contributor to the increasing trend in N<sub>2</sub>O emissions during 1860-2015. Despite the same  
404 input datasets, the interannual variations among the three models were different due to the  
405 differences in model structure and parameters. The estimated N<sub>2</sub>O emissions from VISIT were  
406 consistently higher than those from the other two models during 1860-2015; N<sub>2</sub>O emissions from  
407 DLEM and O-CN were similar in magnitude. The increasing trends of N<sub>2</sub>O emissions before the  
408 end of 1960s were similar among the three models, while the largest increasing trend was found  
409 for O-CN, followed by DLEM and the least for VISIT. The ultimate global terrestrial N<sub>2</sub>O

410 budgets, interannual variations, and attributions of the differences among models will be further  
411 analyzed in more detail after modeling results from all ten models are included.

412         The terrestrial N<sub>2</sub>O emissions showed substantial spatial variations across the global land  
413 surface since 1860 (Figure 6). The highest emission was from the tropical area during all four  
414 periods (i.e., the 1860s, 1900s, 1950s, and 2001-2015) (Figure 6), primarily due to higher soil N  
415 transformation rates and soil N contents in tropical ecosystems. The latitudinal distribution  
416 patterns were slightly different from the 1860s to 2001-2015, showing an increasing importance  
417 and second peak of N<sub>2</sub>O emissions in the temperate climatic zone of the Northern Hemisphere.  
418 Temperate regions were another hotspot for N<sub>2</sub>O emissions due to the high N fertilizer use and N  
419 deposition rates in China, India, Europe, and the contiguous United States. Of all 14 examined  
420 regions as defined by Global Carbon Project (GCP) CH<sub>4</sub> budget synthesis (Saunio et al., 2016),  
421 tropical South America had the largest N<sub>2</sub>O emissions throughout the study period, contributing  
422 to about 20% of the global total emission (Figure 7). China and the contiguous United States  
423 were characterized by the most rapid N<sub>2</sub>O increasing rates. In the recent three decades, China,  
424 India, and western Europe were the only three regions with higher N<sub>2</sub>O emissions from cropland  
425 than that from natural ecosystems. It is noteworthy to point out that the estimated cropland N<sub>2</sub>O  
426 emissions in these three regions have large uncertainty ranges due to varied model representation  
427 and parameterization methods of the impacts from agricultural management. Larger uncertainty  
428 ranges for N<sub>2</sub>O emissions from natural ecosystems were found in Russia, Northern Africa,  
429 Boreal North America, Southeast Asia, and the contiguous United States.

430

## 431 **9. Summary**

432 Current assessments of terrestrial N<sub>2</sub>O emission at regional and global scales are subject  
433 to large uncertainties. The NMIP is attempting to better identify, and eventually reduce those  
434 uncertainties. The activity was initialized in 2015 and currently includes ten terrestrial biosphere  
435 models with N cycling coupled. NMIP is an open initiative and other models are invited to join  
436 the effort. It targets to provide an improved estimate of global and regional terrestrial N<sub>2</sub>O fluxes  
437 as a contribution to the larger GCP Global N<sub>2</sub>O Budget activity. NMIP is being developed with  
438 the capacity to update flux estimates at regular intervals and quantify the uncertainties related to  
439 model structure, algorithms, and parameters. The NMIP protocol includes seven simulation  
440 experiments to quantify and attribute the contribution of environmental factors to the inter-  
441 annual variation and long-term trend of terrestrial N<sub>2</sub>O emissions. In addition, this project  
442 intends to identify our knowledge gaps and bring forward potential strategies for improving the  
443 prediction capability of N<sub>2</sub>O models in the future. The data products and ensemble estimates of  
444 terrestrial N<sub>2</sub>O emissions will be made available and package to be relevant for policy makers  
445 and non-government entities participating in the climate change issues.

446

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461 interest.

462

463 **Appendix A: Brief description of algorithms associated with N<sub>2</sub>O flux in each participating**  
464 **model**

465 1) CLM-CN-N<sub>2</sub>O:

466 CLMCN-N<sub>2</sub>O is based on the DeNitrification-DeComposition (DNDC) model (Li et al.,  
467 1992) implemented in the Community Land Model v3.5 (Oleson et al., 2008; Stöckli et al., 2008)  
468 with explicit carbon and nitrogen (CN) processes (Thornton et al., 2007; Randerson et al., 2009;  
469 Thornton et al. 2009). CLMCN-N<sub>2</sub>O is added to CLM-CN v3.5 in a one-way coupling  
470 framework and simulates N<sub>2</sub>O emissions during nitrification and denitrification processes at an  
471 hourly time step.

472 Nitrification ( $R_{nit}$ ) is temperature and moisture dependent and N<sub>2</sub>O is computed by the  
473 following equation as described in Li et al. (1992):

$$474 R_{nit} = C_{NH_4} f(T1) \tag{1}$$

475 where  $C_{NH_4}$  is the NH<sub>4</sub><sup>+</sup>-N content in soil and  $f(T1)$  is the response function of soil temperature  
476 to nitrification rate.

477 Denitrification is also soil temperature and moisture dependent and it takes place under  
 478 the anaerobic state. CLMCN-N<sub>2</sub>O specifies the anaerobic state when the water-filled pore space  
 479 is more than 41.5% in the soil layer. Under this condition, N<sub>2</sub>O is created based on the growth  
 480 rate of denitrifying bacteria, as well as consumption and assimilation by plants and microbes,  
 481 following Li et al. (1992). Detailed processes in simulating N<sub>2</sub>O emissions can be found in  
 482 Saikawa et al., 2013.

483

484 2) DLEM2.0:

485 The nitrogen cycle scheme in DLEM2.0 (Yang et al. 2015; Xu et al. 2017) are similar as  
 486 DLEM1.0 (Tian et al., 2010, 2011, 2012; Lu and Tian., 2013; Xu et al., 2012), However, the  
 487 N<sub>2</sub>O emission schemes in DLEM2.0 (Xu et al., 2017) have been modified based on Chatskikh et  
 488 al. (2005) and Heinen (2006).

$$489 \quad R_{nit} = k_{nit\_max} f(T1) f(WFPS) C_{NH4} \quad (2)$$

$$490 \quad R_{den} = k_{den\_max} f(T2) f(WFPS) C_{NO3} \quad (3)$$

491 where  $R_{nit}$  is the daily nitrification rate (g N/m<sup>2</sup>/d);  $R_{den}$  is the daily denitrification rate (g  
 492 N/m<sup>2</sup>/d);  $f(T1)$  and  $f(T2)$  are the impact function of daily soil temperature on nitrification and  
 493 denitrification, respectively;  $f(WFPS)$  is the impact function of water-filled pore space (WFPS)  
 494 on nitrification, denitrification and N<sub>2</sub>O diffusion;  $k_{nit\_max}$  is the maximum fraction of NH<sub>4</sub><sup>+</sup>-N  
 495 that is converted to NO<sub>3</sub><sup>-</sup>-N or gases (0-1);  $k_{den\_max}$  is the maximum fraction of NO<sub>3</sub><sup>-</sup>-N that is  
 496 converted to gases (0-1);  $C_{NH4}$  and  $C_{NO3}$  are the soil NH<sub>4</sub><sup>+</sup>-N and NO<sub>3</sub><sup>-</sup>-N content (g N/m<sup>2</sup>). N<sub>2</sub>O  
 497 from denitrification and nitrification processes are calculated as,

$$498 \quad R_{N2O} = (R_{nit} + R_{den}) f(T3) (1 - f(WFPS)) \quad (4)$$

499 where  $R_{N_2O}$  is the daily  $N_2O$  emission rate ( $g\ N/m^2/d$ );  $f(T3)$  is the impact function of daily soil  
 500 temperature on  $N_2O$  diffusion rate from soil pores. The calculation methods for these functions  
 501 and parameters were described in detail in Xu et al. (2017) and Yang et al. (2015).

502

503 3) LM3V-N:

504 In LM3V-N, nitrification is proportional to substrate availability (i.e.,  $NH_4^+$ ), modified by  
 505 functions that account for effects of temperature and WFPS adapted from Parton et al. (1996).

506 Nitrification-associated  $N_2O$  emission ( $R_{nit}$ ) is evaluated by

$$507 R_{nit} = k_{nit\_base} f(WFPS) f(T1) C_{NH_4} / b_{NH_4} \quad (5)$$

508 where  $k_{nit\_base}$  is the base nitrification rate;  $b_{NH_4}$  is the buffer parameter for soil  $NH_4^+$ .

509 Denitrification is described by a Monod-type equation, where both carbon and nitrate  
 510 substrate availability can have limiting effects on N gas production following Li et al. (2000).  
 511 These functions are further modified by temperature (based on Xu-Ri and Prentice, 2008), and  
 512 by WFPS indicating the availability and/or absence of oxygen (adapted from Parton et al., 1996).

$$513 R_{den} = k_{den\_base} f(T2) f(WFPS) f_g C_{NO_3} / b_{NO_3} \quad (6)$$

514 where  $k_{den\_base}$  is the base denitrification rate;  $f_g$  denotes the impact of labile carbon availability  
 515 to nitrate on the growth of denitrifies;  $b_{NO_3}$  is the buffer parameter for soil  $NO_3^-$ .

516 Gaseous losses are partitioned between  $NO_x$  and  $N_2O$  during nitrification is  
 517 parameterized based on air-filled porosity, following Parton et al. (2001). Partitioning between  
 518  $N_2O$  and  $N_2$  during denitrification follows the empirical function of Del Grosso et al. (2000),  
 519 which combines effects of substrate, electron donors (labile C), and water filled pore space.

$$520 R_{N_2O} = 0.004 R_{nit} + R_{den} f(WFPS) f(C_{NO_3}) \quad (7)$$

521 Nitrification/denitrification are treated as fast processes (Shevliakova et al., 2009) and  
522 thus updated on sub-hourly time steps along with updates on soil moisture, soil temperature and  
523 C and N mineralization. Model description including model formulation are detailed in Huang  
524 and Gerber (2015).

525

#### 526 4) LPJ-GUESS:

527 The nitrogen cycle scheme in LPJ-GUESS is based on CENTURY (Parton et al., 1996)  
528 and Xu-Ri and Prentice (2008). Inorganic soil nitrogen pools in the model are: ammonium,  
529 nitrite and nitrate. Nitrification only occurs in the dry part of the soil (fractionated using water-  
530 filled pore space, WFPS), the ratio between N<sub>2</sub>O and NO<sub>x</sub> of the gaseous losses in nitrification is  
531 based on the moisture content in the soil (f(WFPS)).

$$532 R_{nit} = k_{nit\_max} f(WFPS) C_{NH4} \quad (8)$$

533 Denitrification occurs in the wet part (based on WFPS) of the soil, the denitrification rate  
534 depends on temperature, soil moisture and labile carbon (approximated with heterotrophic  
535 respiration,  $rh$ ). Gaseous losses through denitrification results in N<sub>2</sub>O, N<sub>2</sub> and NO<sub>x</sub>.

$$536 R_{den} = k_{den\_base} f(T2) f(WFPS) f(rh) C_{NO3} \quad (9)$$

537 The fractionation between the gaseous N species are modelled using soil moisture and  
538 temperature. All losses of gaseous N, are modelled. Emissions to the atmosphere from these  
539 pools are modelled using rate modifiers that are based on the soil moisture and temperature. No  
540 re-transformation of these gaseous N species is considered. These processes (N-cycling and  
541 gaseous N emissions) are modelled in different land use classes: natural vegetation,  
542 pastures/rangelands and croplands. On croplands, fertilizers are spread as mineral and/or organic  
543 N. Mineral fertilizers are considered as an input to the ammonium and nitrate pools at a fixed



544 ratio (50/50), and manure as an input into the organic nitrogen pool with a fixed C:N ratio  
545 (currently set to 30).

546

547 5) LPX-Bern:

548 The implementation of nitrogen dynamics in LPX-Bern is based on the work of Xu-Ri  
549 and Prentice (2008). Nitrogen uptake by plants is governed by their demand and the availability  
550 of nitrogen in two soil pools representing ammonium and nitrate. Nitrogen from deposition and  
551 fertilization are added to these inorganic soil pools. Losses include ammonium volatilization,  
552 nitrate leaching as well as N<sub>2</sub>O and NO production during nitrification and N<sub>2</sub>O, NO and N<sub>2</sub>  
553 production during denitrification. Aerobic nitrification of ammonium is dependent on soil  
554 temperature ( $T_{soil}$ ) and indirectly on soil water content due to the partitioning of wet and dry soil:

$$555 R_{nit} = max_{nit} f_1(T_{soil}) C_{NH_4, dry} \quad (10)$$

556 where  $max_{nit} = 0.92 \text{ day}^{-1}$  is the daily maximum nitrification rate at 20°C.

557 Anaerobic denitrification of nitrate in wet soil depends on labile carbon availability and soil  
558 temperature:

$$559 R_{den} = R_{mb} / (R_{mb} + K_{mb}) f_2(T_{soil}) C_{NO_3, wet} / (C_{NO_3, wet} + K_n) \quad (11)$$

560 The parameters  $K_{mb}$  and  $K_n$  are taken from Xu-Ri and Prentice (2008) and  $R_{mb}$  is the  
561 microbiological soil respiration. The amount of Nitrogen lost as N<sub>2</sub>O due to nitrification and  
562 denitrification is modelled as a function of soil temperature, water content and the respective  
563 process rate.

564

565 6) O-CN:

566 The treatment of inorganic soil nitrogen dynamics in O-CN follows largely Xu-Ri and

567 Prentice (2008). O-CN (Zaehle and Friend, 2010) considers N losses to NH<sub>3</sub> volatilisation, NO<sub>x</sub>,  
 568 N<sub>2</sub>O and N<sub>2</sub> production and emission, as well as NH<sub>4</sub> and NO<sub>3</sub> leaching. Inorganic nitrogen  
 569 dynamics in the soil are tightly coupled to plant uptake and net mineralization. The anaerobic  
 570 volume fraction of the soil is estimated by an empirical function of the fractional soil moisture  
 571 content (Zaehle et al. 2011). The fraction of ammonium in the aerobic part of the soil is subject  
 572 to nitrification, according to:

$$573 \quad R_{nit} = v_{max_{nit}} f(T1) f(pH1) C_{NH4} \quad (12)$$

574 where  $f(pH1)$  is the soil pH response functions for nitrification (Li et al. 1992; Xu-Ri and  
 575 Prentice, 2008), and  $v_{max_{nit}}$  is the maximum daily nitrification rate under 20°C and favourable  
 576 pH conditions (Xu-Ri and Prentice, 2008).

577 Gross denitrification of the fraction of nitrate under anoxic conditions is modelled as:

$$578 \quad R_{den} = R_{mb} / (R_{mb} + K_{mb}) f(T2) f(pH2) C_{NO3} / (C_{NO3} + K_n) \quad (13)$$

579 where  $f(pH2)$  is the soil pH response functions for denitrification (Li et al., 1992; Xu-Ri and  
 580 Prentice, 2008),  $R_{mb}$  is the soil microbial respiration rate, and  $K_{mb}$  and  $K_n$  parameters taken from  
 581 Li et al. (1992).

582 The N<sub>2</sub>O production from nitrification and denitrification is then calculated as:

$$583 \quad R_{N2O} = a_{nit} f(T1) R_{nit} + b_{den} f(T2) f(pH3) R_{den} \quad (14)$$

584 where  $a_{nit}$  and  $b_{den}$  are fraction loss constants,  $f(pH3)$  is a pH-modifier changing the degree of  
 585 denitrification producing N<sub>2</sub>O versus NO<sub>x</sub> or N<sub>2</sub> (Zaehle et al. 2011). Emissions of volatile  
 586 compounds are simulated using the empirical emission of Xu-Ri and Prentice (2008).

587

588 7) ORCHIDEE:

589 Modeling of the mineral N dynamics by the ORCHIDEE model originates from the  
 590 formulations used in the O-CN (Zaehle and Friend, 2010). It is composed of five pools for  
 591 ammonium/ammoniac, nitrate, NO<sub>x</sub>, nitrous oxide, and di-nitrogen forms. N<sub>2</sub>O production in  
 592 both nitrification and denitrification processes are represented.

593 The potential daily rate of nitrification,  $R_{nit}$ , occurs only on the aerobic fraction of the  
 594 soil and is a function of temperature, pH, and ammonium concentration ( $C_{NH4}$ ):

$$595 R_{nit} = (1 - f(WFPS))f(T1)f(pH1)k_{nit}C_{NH4} \quad (15)$$

596 where  $k_{nit}$  is the reference potential NO<sub>3</sub><sup>-</sup> production per mass unit of ammonium.

597 N<sub>2</sub>O production by nitrification ( $R_{N2O,nit}$ , g N-N<sub>2</sub>O/m<sup>2</sup>/d) is expressed as a function of the  
 598 potential daily rate of nitrification ( $R_{nit}$ , g N-NO<sub>3</sub><sup>-</sup>/m<sup>2</sup>/d), temperature and the water content as  
 599 shown in Zhang et al. (2002).

$$600 R_{N2O,nit} = f(WFPS)f(T1)R_{nit}p_{N2O,nit} \quad (16)$$

601 where  $p_{N2O,nit}$  (g N-N<sub>2</sub>O (g N-NO<sub>3</sub><sup>-</sup>)<sup>-1</sup>) is the reference N<sub>2</sub>O production per mass unit of NO<sub>3</sub><sup>-</sup>  
 602 produced by nitrification. The denitrification occurs on the anaerobic fraction of the soil which is  
 603 computed as a function of the water-filled porosity ( $f(WFPS)$ ) and is controlled by temperature,  
 604 pH, soil NO concentration and denitrifier microbial activity ( $a_{microb}$ , g m<sup>-2</sup>) (Li et al., 2000).

$$605 R_{N2O,den} = f(WFPS)f(T2)f(pH)f(NO)p_{N2O,den}a_{microb} \quad (17)$$

606 where  $f(NO)$  is a Michaelis-Menten shape function and  $p_{N2O,den}$  is the reference N<sub>2</sub>O production  
 607 per mass unit of denitrifier microbes.

608

## 609 8) ORCHIDEE-CNP

610 ORCHIDEE-CNP (Goll et al., 2017) is a version with the implementation of the  
 611 phosphorus cycle into the nitrogen enabled version of ORCHIDEE (ORCHIDEE-CN; Vuichard

612 in prep.). The inorganic soil nitrogen dynamics of ORCHIDEE-CNP includes N<sub>2</sub>O from both  
613 nitrification and denitrification processes following the processes of O-CN (Zaehle et al., 2011).  
614 One exception is the BNF. In ORCHIDEE-CNP, BNF is a function of NPP (Cleveland et al.,  
615 1999) and also regulated by soil mineral N concentration. ORCHIDEE-CNP accounts for  
616 influence of phosphorus state of vegetation on tissue nutrient concentrations and phosphatase-  
617 mediated biochemical mineralization. Changes in nutrient content (quality) of litter affect the  
618 carbon use efficiency of decomposition and in return the nutrient availability to vegetation. The  
619 model explicitly accounts for root zone depletion of phosphorus as a function of root phosphorus  
620 uptake and phosphorus transport from soil to the root surface.

621

## 622 9) TRIPLEX-GHG

623 TRIPLEX-GHG model (Zhu et al., 2014; Zhang et al., 2017) is designed to simulate N<sub>2</sub>O  
624 emissions by coupling major theoretical foundations for processes of nitrification and  
625 denitrification reported by Li et al. (2000). Briefly, the nitrification rate is calculated by the  
626 Michaelis-Menten function based on the concentration of NH<sub>4</sub><sup>+</sup>, and microbial activity of  
627 nitrifying bacteria is explicitly involved based on simulating their growth and death;  
628 denitrification is expressed in a more complex way by taking into account the chain reaction  
629 ( $NO_3^- \rightarrow NO_2^- \rightarrow NO \rightarrow N_2O \rightarrow N_2$ ). Each step of denitrification can be regarded as an  
630 independent process, but these steps are linked by competition for DOC between specific  
631 denitrifiers during each step. A double substrate-based (DOC and NO<sub>x</sub>) Michaelis-Menten  
632 equation was adopted to simulate the growth rates of NO<sub>x</sub> denitrifiers (Li et al., 2000). In  
633 addition, the effects of different factors, such as soil temperature, soil moisture and pH, are also  
634 considered. The key equations for nitrification are showed as follows,

635  $R_{nit} = B_{nit} \frac{R_{max} C_{NH_4}}{(6.18 + C_{NH_4})} pH$  (18)

636  $R_{max} = COE_{NR} \cdot N_p$  (19)

637  $F_{NN_2O} = FMAX_{N_2O} R_{nit} f(T1) f(WFPS)$  (20)

638 where  $R_{nit}$  is the nitrification rate ( $\text{kg N m}^{-2} \text{d}^{-1}$ );  $R_{max}$  is the maximum nitrification rate ( $\text{d}^{-1}$ );  $B_{nit}$   
 639 is the biomass concentration of nitrifiers ( $\text{kg C m}^{-2}$ );  $pH$  is the soil pH;  $COE_{NR}$  represents the  
 640 nitrification coefficient;  $N_p$  represents the nitrification potential ( $\text{mg N kg}^{-1} \text{d}^{-1}$ );  $FMAX_{N_2O}$  is the  
 641 maximum  $\text{N}_2\text{O}$  fraction during nitrification ( $\text{kg N m}^{-2} \text{d}^{-1}$ ); and  $f(T1)$  and  $f(WFPS)$  are the  
 642 functions of the effects of soil temperature and soil moisture on  $\text{N}_2\text{O}$  emissions during  
 643 nitrification, respectively.

644 The key equations for denitrification are showed as follows,

645  $R_{NO_x} = MUE_{NO_x} \frac{[DOC]}{[DOC] + K_c} \frac{[NO_x]}{[NO_x] + K_n}$  (21)

646  $F_{ANNO_x} = COE_{dNO_x} B_{denit} \left( \frac{R_{NO_x}}{EFF_{NO_x}} + \frac{MAI_{NO_x} [NO_x]}{[N]} \right) f_{NO_x}(pH2) f(T2)$  (22)

647 where  $MUE_{NO_x}$  is the maximum growth rate of  $\text{NO}_x$  denitrifiers ( $\text{h}^{-1}$ );  $[DOC]$  and  $[NO_x]$  represent  
 648 the concentrations of DOC ( $\text{kg C m}^{-3} \text{h}^{-1}$ ) and  $\text{NO}_x$  ( $\text{kg N m}^{-3} \text{h}^{-1}$ ), respectively, in the anaerobic  
 649 balloon;  $K_c$  ( $\text{kg C m}^{-3}$ ) and  $K_n$  ( $\text{kg N m}^{-3}$ ) are the half saturation value of C and N oxides,  
 650 respectively.  $F_{ANNO_x}$  is the consumption rate of  $\text{NO}_x$  ( $\text{kg N m}^{-3} \text{h}^{-1}$ );  $COE_{dNO_x}$  represents the  
 651 coefficient of  $\text{NO}_x$  consumption;  $B_{denit}$  is the biomass of denitrifiers ( $\text{kg C m}^{-3}$ );  $R_{NO_x}$  is the  $\text{NO}_x$   
 652 reduction rate ( $\text{h}^{-1}$ );  $[NO_x]$  and  $[N]$  are the concentrations of  $\text{NO}_x$  and total N, respectively, in the  
 653 anaerobic balloon ( $\text{kg N m}^{-3}$ );  $EFF_{NO_x}$  is the efficiency parameter for  $\text{NO}_x$  denitrifiers ( $\text{kg C kg N}$   
 654  $^{-1}$ );  $MAI_{NO_x}$  is the maintenance coefficient of  $\text{NO}_x$  ( $\text{h}^{-1}$ ); and  $f(t)_{denit}$  represents the effect of the  
 655 soil temperature on the denitrification rate during each step.

656

657 10) VISIT

658 The nitrogen cycle scheme of VISIT is composed of three organic soil nitrogen pools  
659 (microbe, litter, and humus), two inorganic soil nitrogen pools (ammonium and nitrate), and  
660 vegetation pools. Fertilizer is considered as an input to the ammonium and nitrate pools at a  
661 fixed ratio, and manure as an input into the litter organic nitrogen pool. N<sub>2</sub>O emissions through  
662 nitrification and denitrification are estimated using the scheme developed by Parton et al. (1996).

663 Nitrification-associated N<sub>2</sub>O emission ( $R_{nit,N_2O}$ ) is evaluated as follows,

$$664 R_{nit,N_2O} = f(WFPS)f(pH1)f(T1)(K_{max} + F_{max}f(NH_4)) \quad (23)$$

665 where  $K_{max}$  is the soil-specific turnover coefficient;  $F_{max}$  is the parameter of maximum  
666 nitrification gas flux; and  $f(NH_4)$  is the effect of soil ammonium on nitrification. Denitrification-  
667 associated N<sub>2</sub>O emission ( $R_{den,N_2O}$ ) is evaluated by the following equation:

$$668 R_{den,N_2O} = R_{den}(1 + R_{N_2/N_2O}) \quad (24)$$

$$669 R_{den} = \min(f(NO_3), f(CO_2)) \times f(WFPS) \quad (25)$$

670 where  $R_{N_2/N_2O}$  is the fractionation coefficient, which is also a function of WFPS, soil nitrate, and  
671 heterotrophic respiration,  $f(NO_3)$  is the maximum denitrification rate in high soil respiration  
672 rate condition,  $f(CO_2)$  is the maximum denitrification rate in high NO<sub>3</sub><sup>-</sup> levels, and  $f(WFPS)$  is  
673 the effect of WFPS on denitrification rate.

674

675

676 **References**

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949 **Figure Captions**

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980 DLEM, VISIT, and O-CN model simulations. The error bars denote ± 1 standard deviation.

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982 **Table 1.** Summary of the NMIP driving forces  
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<b>Data name</b>	<b>Period</b>	<b>Temporal resolution</b>	<b>Spatial resolution</b>	<b>Sources</b>	<b>Variables</b>
Climate	1901-2015	6-hourly	0.5°	CRU-NCEP	Incoming longwave / shortwave radiation, air humidity, pressure, precipitation, temperature, and wind speed
CO <sub>2</sub>	1860-2015	Monthly	0.5°	NCAR	CO <sub>2</sub> concentration
N deposition	1860-2015	Yearly	0.5°	Eyring et al. (2013)	NH <sub>x</sub> -N and NO <sub>y</sub> -N deposition
N Fertilizer use	1860-2014	Yearly	0.5°	Lu and Tian (2017)	N fertilizer use rate in cropland
Manure N input	1860-2014	Yearly	0.5°	Zhang et al. (2017)	Manure N production
Potential vegetation	One time	One time	0.5°	SYNMAP	Fraction of natural vegetation types
Cropland	1860-2015	Yearly	0.5°	HYDE 3.2	Cropland fraction

984 Note: Detailed descriptions of the major NMIP model input datasets have been provided in  
 985 previous publications or online documents. Here we only provide a brief description of sources  
 986 and spatiotemporal patterns of these datasets.

987 **Table 2.** Participating models

<b>Model</b>	<b>Full name</b>	<b>Contact</b>	<b>Affiliation</b>	<b>Citation</b>
CLM-CN	Community land model - CN	Eri Saikawa	Emory University, USA	Saikawa et al. (2013)
DLEM	Dynamic Land Ecosystem Model	Hanqin Tian	Auburn University	Tian et al. (2015) Xu et al. (2017)
LM3V-N	Land Model 3V-N	Stefan Gerber	University of Florida	Huang and Gerber (2015)
LPJ-GUESS	Lund-Potsdam-Jena General Ecosystem Simulator	Stefan Olin/ Almut Arneth	Lund University, Sweden/KIM, Dept. Atmospheric Environmental Research, Germany	Olin et al. (2015); Xu-Ri and Prentice (2008)
LPX-Bern	Land Processes and eXchanges model - Bern	Sebastian Lienert/ Fortunat Joos	Institute for Climate and Environmental Physics, University of Bern, Switzerland	Stocker et. al. (2013) Xu-Ri and Prentice (2008)
O-CN	ORCHIDEE-CN	Sönke Zaehle	Max Planck Institute for Biogeochemistry	Zaehle et al. (2011)
ORCHIDEE	Organising Carbon and Hydrology In Dynamic Ecosystems	Nicolas Vuichard	IPSL – LSCE, France	Vuichard et al. (in prep)
ORCHIDEE-CNP	Organising Carbon and Hydrology In Dynamic Ecosystems-CNP	Jinfeng Chang/ Daniel Goll	IPSL – LSCE, France	Goll et al., 2017
TRIPLEX-GHG	TRIPLEX-GHG	Changhui Peng	University of Quebec at Montreal, Canada	Zhu et al. (2014); Zhang et al. (2017)
VISIT	Vegetation Integrated Simulator for Trace-gases	Akihiko Ito	National Institute for Environmental Studies, Japan	Inatomi et al. (2010); Ito and Inatomi (2012)

**Table 3.** Model characteristics in simulating major N cycling processes

	CLM-CN	DLEM	LM3V-N	LPJ -GUESS	LPX-Bern	O-CN	ORCHIDEE	ORCHIDEE -CNP	TRIPLEX-GHG	VISIT
Open N cycle <sup>1</sup>	yes	yes	yes	yes	yes	yes	yes	yes	yes	yes
C-N coupling	yes	yes	yes	yes	yes	yes	yes	yes	yes	yes
N pools <sup>2</sup>	(13, 3, 4)	(6, 6, 8)	(6, 4, 3)	(5, 6, 11)	(4,3,8)	(9, 6, 9)	(9, 6, 9)	(9, 6, 9)	(3, 9, 4)	(4, 1, 4)
Demand and supply-driven Plant N uptake	yes	yes	yes	yes	yes	yes	yes	yes	yes	yes
N allocation <sup>3</sup>	Dynamic	Dynamic	Dynamic	Dynamic	Dynamic	Dynamic	Dynamic	Dynamic	Dynamic	Dynamic
Nitrification	$f(T, SWC)$	$f(T, SWC, C_{NH4})$	$f(T, SWC, C_{NH4})$	$f(T, SWC, C_{NH4})$	$f(T, SWC, C_{NH4})$	$f(T, SWC, pH, C_{NH4})$	$f(T, SWC, pH, C_{NH4})$	$f(T, SWC, pH, C_{NH4})$	$f(pH, C_{NH4}, T, SWC)$	$f(T, SWC, pH, C_{NH4})$
Denitrification	$f(T, SWC, C_{NO3})$	$f(T, clay, Rh, C_{NO3})$	$f(T_{soil}, Rh, SWC, C_{NH4}, C_{NO3})$	$f(T, Rh, SWC, C_{NO3})$	$f(T, SWC, R_{mb}, C_{NO3})$	$f(T, SWC, pH, R_{mb}, C_{NO3})$	$f(T, SWC, pH, denitrifier, C_{NO3})$	$f(T, SWC, pH, R_{mb}, C_{NO3})$	$f(DOC, C_{NO3}, pH, T_{soil})$	$f(SWC, Rh, C_{NO3})$
Mineralization, immobilization	$f(C:N)$	$f(C:N)$	$f(C_{NO3}, C_{NH4})$	$f(C:N)$	$f(C:N)$	$f(C:N)$	$f(C:N)$	$f(C:N)$	$f(C:N)$	$f(C:N)$
N leaching	$f(runoff)$	$f(runoff)$	$f(runoff)$	$f(runoff)$	$f(runoff)$	$f(runoff, clay)$	$f(runoff)$	$f(runoff)$	$f(runoff)$	$f(runoff)$
NH <sub>3</sub> volatilization	$f(C_{NH4})$	$f(pH, T, SWC, C_{NH4})$	$f(pH, T, SWC, C_{NH4})$	$f(pH, T, SWC, C_{NH4})$	$f(pH, T, SWC, C_{NH4})$	$f(pH, C_{NH4})$	$f(pH, C_{NH4})$	$f(pH, C_{NH4})$	$f(pH, C_{NH4})$	$f(pH, T, SWC, C_{NH4})$
Plant N turnover <sup>4</sup>	Dynamic	Dynamic	Dynamic	Dynamic	Dynamic	Dynamic	Dynamic	Dynamic	Dynamic	Dynamic
N resorption	$f(C:N)$	$f(C:N)$	fixed	Crop: dynamic, the rest: fixed	$f(N_{leaf})$	fixed	$f(N_{leaf})$	fixed	$f(C:N)$	fixed
N fixation	$f(NPP)$	Fixed	$f(C_{NH4}, C_{NO3}, light, plant demand)$	$f(ET)$	<i>Implied by mass balance</i>	$f(C_{cost}, C_{root})$	$f(ET)$	$f(NPP)$	$f(biomass)$	$f(ET)$
N fertilizer use	no	yes	yes	yes	yes	yes	yes	yes	yes	yes
Manure N use	no	yes	yes	yes	no	no	yes	yes	no	yes
N deposition	yes	yes	yes	yes	yes	yes	yes	yes	yes	yes

989 Note: <sup>1</sup> “open” denotes that excess N can be leached from the system; <sup>2</sup> numbers of N pools (vegetation pools, litter pools, soil pools); <sup>3</sup> Dynamic denotes time-varied N allocation  
990 ratio to different N pools; <sup>4</sup> turnover time for various vegetation nitrogen pools. T: soil temperature; Clay: soil clay fraction; ET: evapotranspiration; Biomass: vegetation carbon;  
991 NPP: net primary production; Nleaf: leaf N concentration; Runoff: soil surface and drainage runoff; Ccost: carbon cost during N<sub>2</sub> fixation; SWC: soil water content; denitrifier: soil  
992 denitrifier biomass; Rh: soil heterogeneous respiration.

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**Table 4.** List of nitrogen and carbon variables provided by NMIP models

<b>Name of Variables</b>	<b>Unit</b>	<b>Frequency</b>	<b>Level</b>
<b>Nitrogen Fluxes:</b>			
N <sub>2</sub> O flux, Biological N fixation, Plant N uptake (sum of ammonium and nitrate), Net N mineralization, Nitrification rate, Denitrification rate, N leaching (DIN, DON, PON or total N leaching), NH <sub>3</sub> volatilization	kg N m <sup>-2</sup> s <sup>-1</sup>	Monthly (1980-2015) Annual (1860-2015)	Grid- and biome-levels
<b>Nitrogen Pools:</b>			
N in Vegetation, N in Above-ground Litter Pool, N in Soil (including below-ground litter), N in Products pools	kg N m <sup>-2</sup>	Annual (1860-2015)	Grid-levels
<b>Carbon Fluxes:</b>			
Gross Primary Production, Autotrophic (Plant) Respiration, Net Primary Production, Heterotrophic Respiration	kg C m <sup>-2</sup> s <sup>-1</sup>	Monthly (1980-2015) Annual (1860-2015)	Grid- and biome-levels
<b>Carbon Pools:</b>			
C in Vegetation, C in Above-ground Litter Pool, C in Soil (including below-ground litter), C in Products pools, C in Vegetation	kg C m <sup>-2</sup>	Annual (1860-2015)	Grid-level Biome-level (vegetation C)

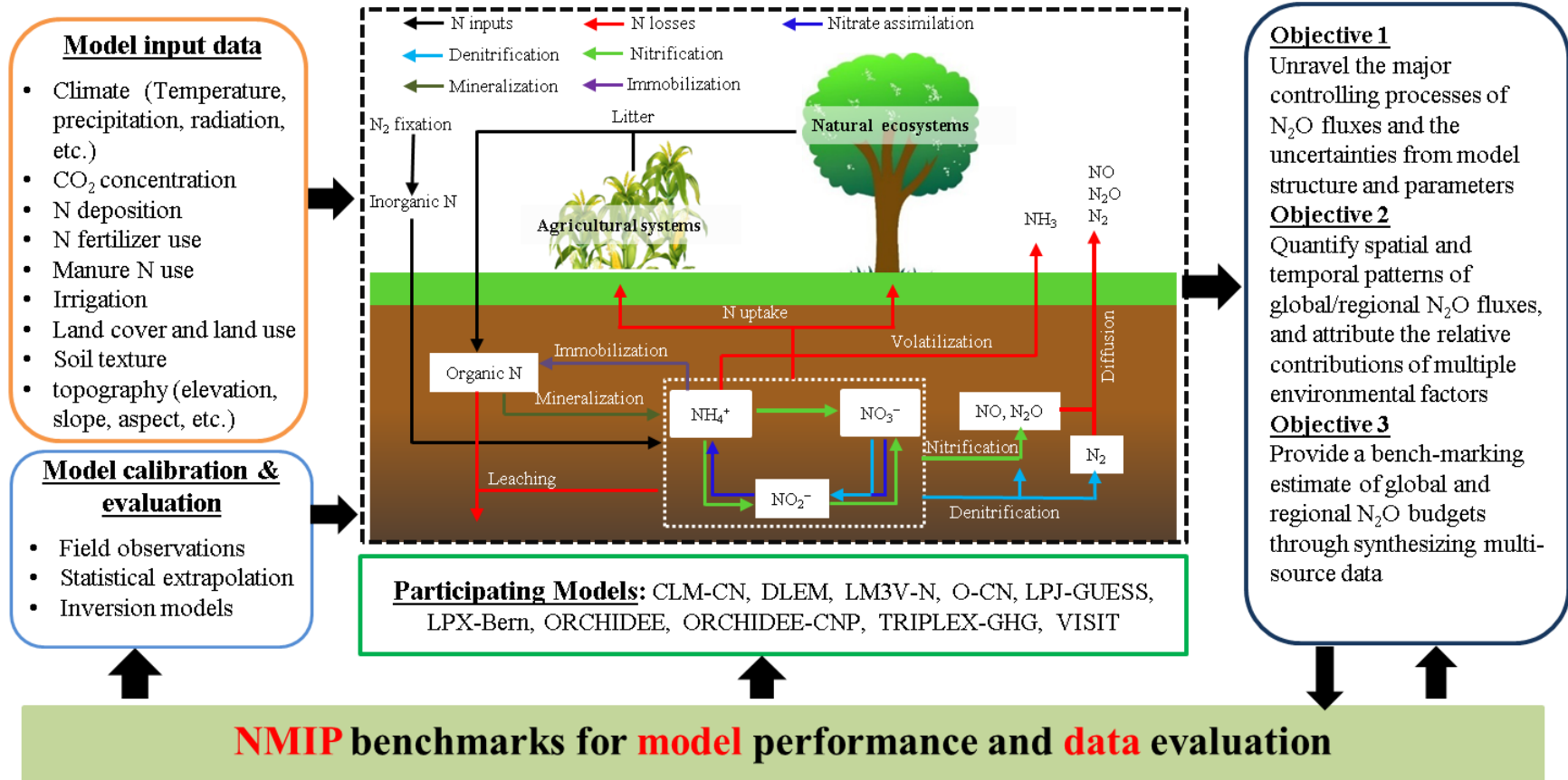
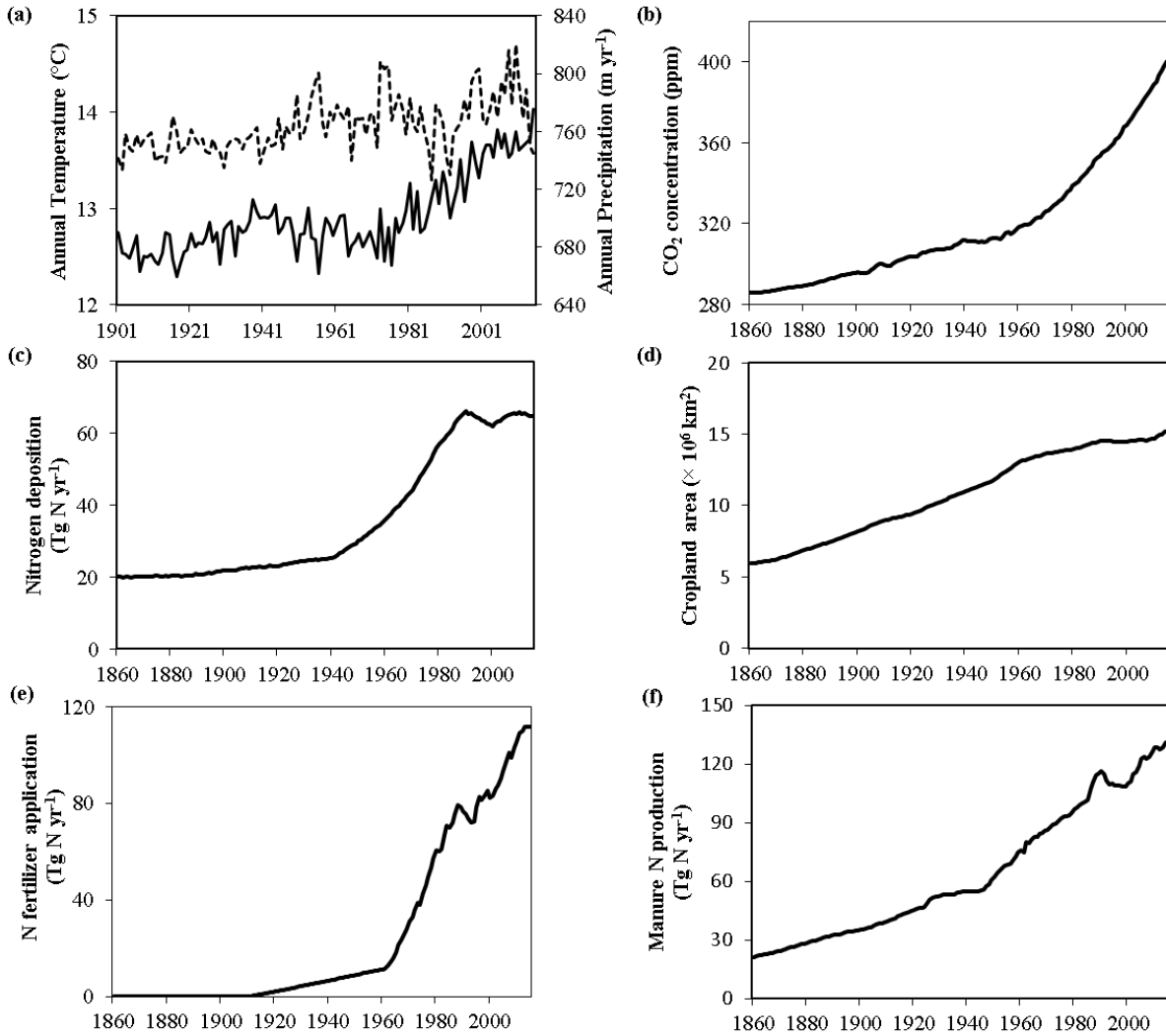
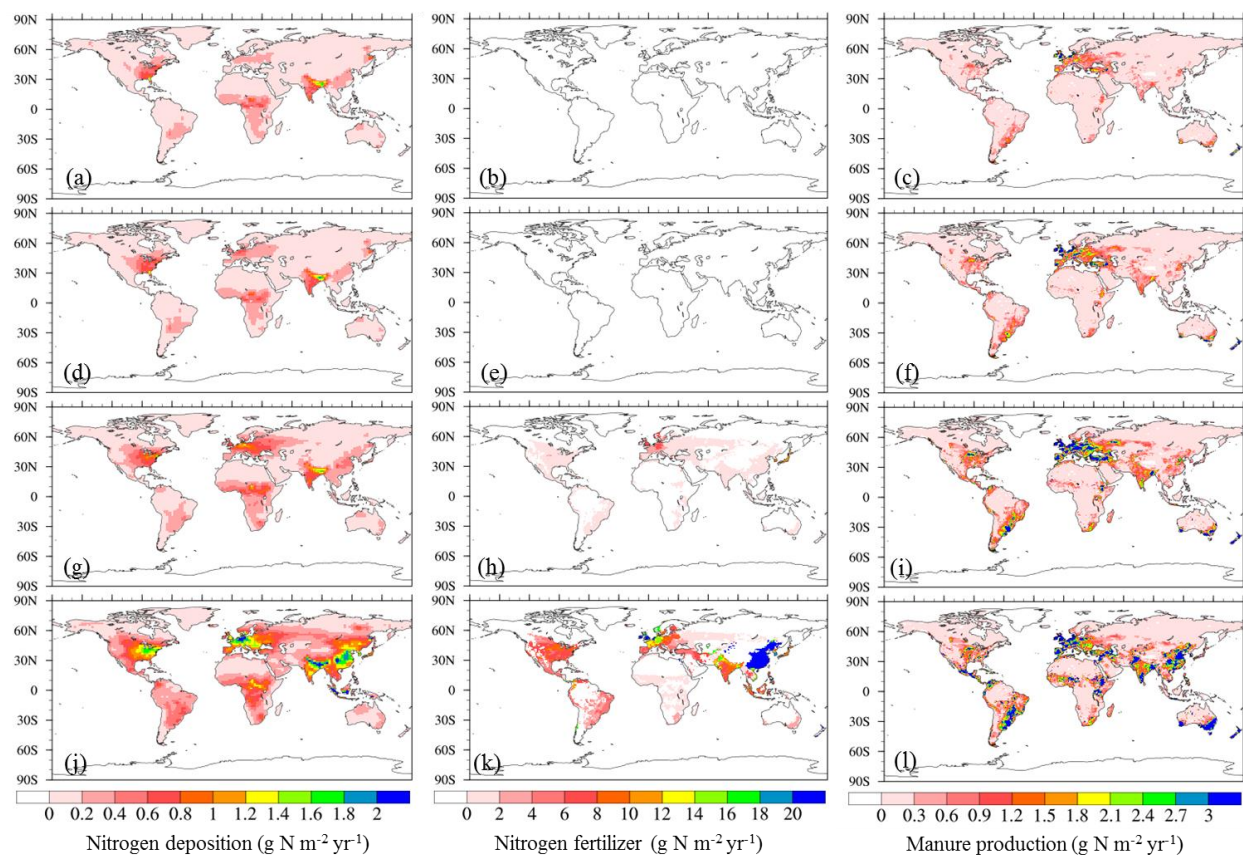


Figure 1. The framework of the N<sub>2</sub>O Model Intercomparison project (NMIP)



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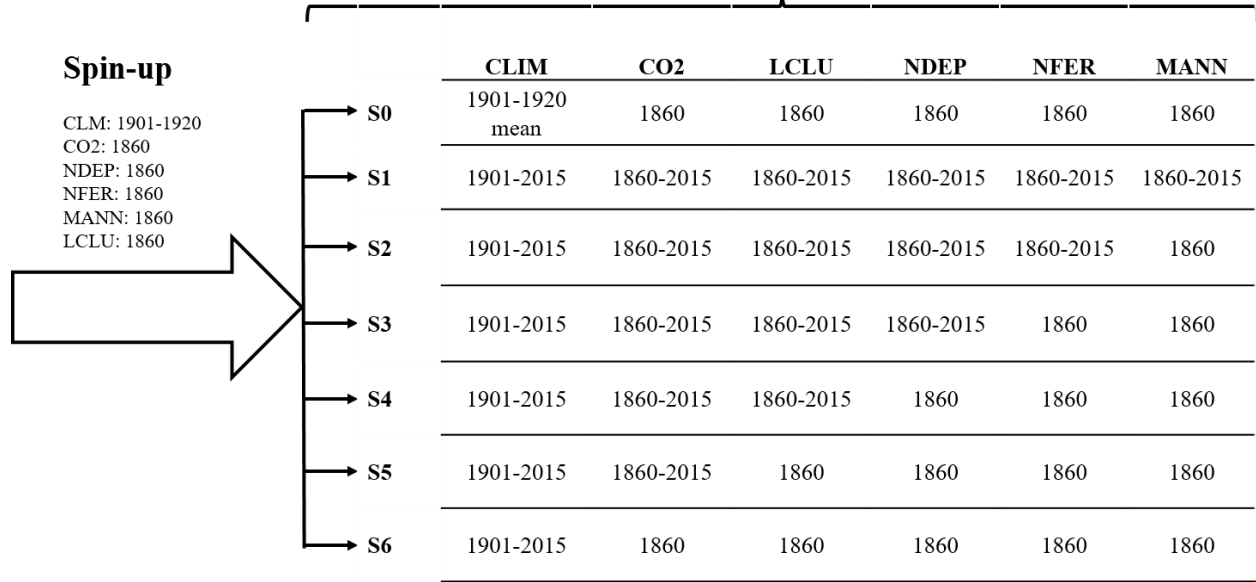
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**Transient Run (1860-2015)**

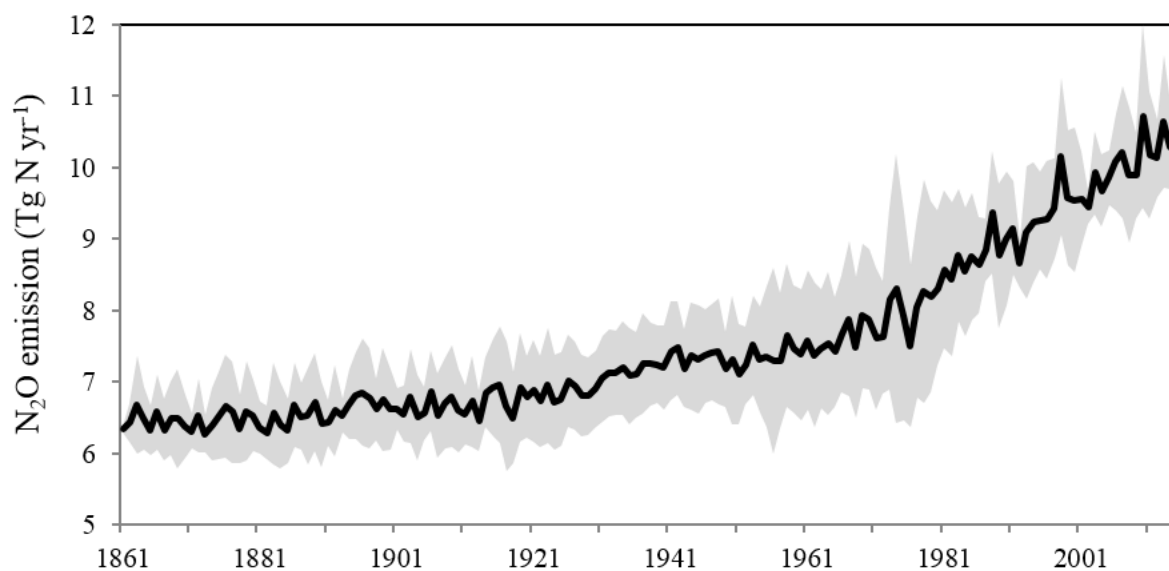


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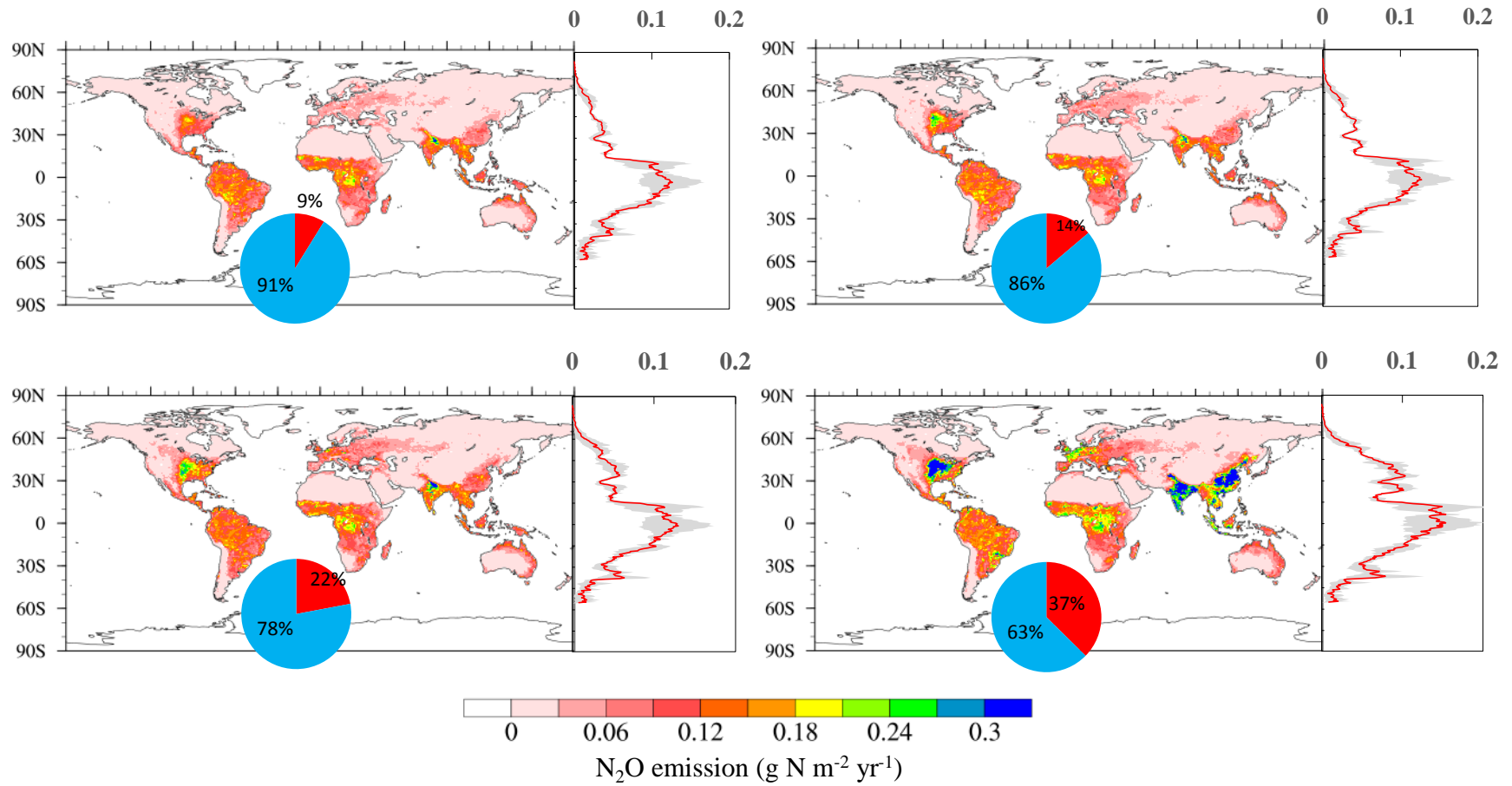
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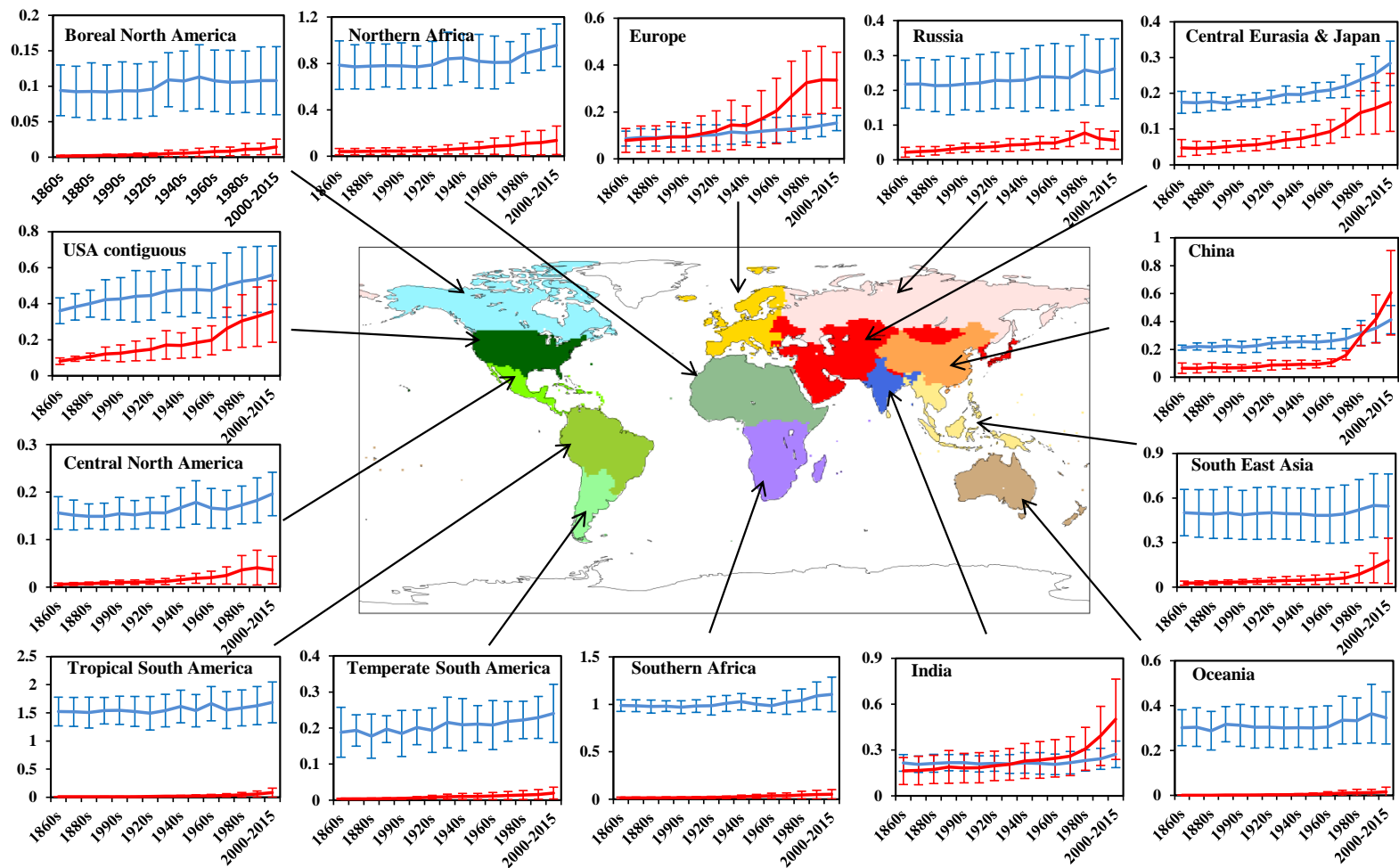
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